

Near-surface Qs and Qp estimations from Rayleigh waves using multi-channel analysis of surface waves (MASW) at an Arctic ice-sheet site

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Summary

We analyze seismic data acquired on the Jakobshavn Glacier, Greenland for reflection surveys. With the use of the multi-channel analysis of surface waves (MASW) seismic method Rayleigh-wave attenuation versus frequency curves were measured and inverted not only for shear-wave quality factor Qs but also for compressional-wave quality factor Qp, which was possible as a result of the favorable relationship between the compressional (P)-wave (Vp) and shear (S)-wave velocity (Vs), i.e., Vp/Vs ratio <2.2 typical for ice sheets. Changes in the a-priori Vp model used for the inversion had little impact on the resulting Qs, but impacted the Qp results, while still preserving the overall Qp trend. This suggests that Qs can be estimated with higher degree of reliability even when Vp errors are assumed. As well, this research showed the benefits of the seismic method for studying glaciers, i.e., that a single seismic survey can provide 3 sets of parameters, Vp reflection results, Vs sections, and Qs and Qp sections.

Introduction

Seismic reflection data were recorded on Jakobshavn Glacier, Greenland (Horgan et al., 2008). Further analysis revealed that these data can be successfully used for shear S-wave velocity (Vs) estimations using surface-wave analysis. This research also showed that acquired data and obtained results can be also used for Rayleigh-wave analysis for evaluating quality factors Q. We focus on multichannel analysis of surface waves (MASW) approach, which was initially used in previous research to evaluate ice-sheet Vs (Ivanov et al., 2009) to obtain quality factor Q estimates.

Stiffness properties of near-surface materials are important for various environmental and engineering applications. Stiffness is directly related to Vs, which increases as material shear strength (stiffness, rigidity) increases. Vs can be estimated by analyzing surface waves on seismic data records. We use the MASW method to estimate near-surface shear-wave velocity from high-frequency (≥ 2 Hz) Rayleigh-wave data (Song et al., 1989; Park et al., 1998; Miller et al., 1999b; Xia et al., 1999). Shear-wave velocities estimated using MASW have been reliably and consistently correlated with drill data. Using the MASW method, (Xia et al., 2000) non-invasively measured Vs within 15% of Vs measured in wells. (Miller et al., 1999b) mapped bedrock with 0.3-m (1-ft) accuracy at depths of about 4.5-9 m (15-30 ft), as confirmed by numerous borings.

The MASW method has been applied to problems such as characterization of pavements (Ryden et al., 2004), the study of Poisson's ratio (Ivanov et al., 2000a), study of levees and subgrade (Ivanov et al., 2004; Ivanov et al., 2006b), investigation of sea-bottom sediment stiffness (Ivanov et al., 2000b; Kaufmann et al., 2005; Park et al., 2005), mapping of fault zones (Ivanov et al., 2006a), study of Arctic ice sheets (Tsoflias et al., 2008; Ivanov et al., 2009), detection of dissolution features (Miller et al., 1999a), and measurement of Vs as a function of depth (Xia et al., 1999). Applications of the MASW method have been extended to determination of near-surface quality factor Q (Xia et al., 2013) and the acquisition of more realistic compressional-wave refraction models (Ivanov et al., 2006c; Ivanov et al., 2010; Piatti et al., 2013). A review of established approaches of surface wave methods (SWM) can be found in Socco et al. (2010). Most recent developments of the SWM include the expansion with the use of the horizontal component of the Rayleigh wave (Boaga et al., 2013), the simultaneous use of guided-waves with multi-mode surface waves in land and shallow marine environments (Boiero et al., 2013) and evaluation at landfill sites (Suto, 2013).

Shear-wave quality factor, Qs, (dissipation factor Q_s^{-1} , a.k.a. intrinsic attenuation) and conditionally compressional-wave quality factor, Qp (dissipation factor Q_p^{-1}), can be estimated from the MASW seismic data. Rayleigh-wave attenuation-frequency curves can be measured (Xia et al., 2012; Xia et al., 2013) and then using estimated MASW Vs with available compressional P-wave (Vp) and density information numerical relationship between these variables can be applied Anderson et al. (1965).

$$\alpha_R(f) = \frac{\pi f}{c_R^2(f)} \left[\sum_{i=1}^n P_i(f) Q_{P_i}^{-1} + \sum_{i=1}^n S_i(f) Q_{S_i}^{-1} \right] \quad (1)$$

Where $P_i(f) = V_{P_i} \frac{\partial c_R(f)}{\partial V_{P_i}}$, $S_i(f) = V_{S_i} \frac{\partial c_R(f)}{\partial V_{S_i}}$. $\alpha_R(f)$ is the Rayleigh-wave attenuation coefficient in 1/length, f is frequency in Hz, Q_{P_i} and Q_{S_i} are the quality factors for P and S waves of the i^{th} layer, V_{P_i} and V_{S_i} are the P- and S-wave velocities of the i^{th} layer. The suggested inversion system is sensitive to Qp only when Vp/Vs ratio is smaller than 2.2, (i.e., Vs/Vp > 0.45) (Xia et al., 2002). Attenuation versus frequency curves can be estimated in the frequency domain by measuring amplitude attenuation trend at a given frequency across a whole record (Xia et al., 2012).

The MASW method is applied by performing the following steps. A single seismic-data record is acquired. These data are

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transformed into a dispersion-curve image (Park et al., 1998; Luo et al., 2009), which is used to evaluate a dispersion-curve trend(s) of the Rayleigh wave. This curve is then inverted to produce a 1D Vs model (Xia et al., 1999). By assembling numerous 1D Vs models, derived from consecutive seismic shot records, 2D (Miller et al., 1999b) or 3D (Miller et al., 2003) Vs models can be obtained.

Inverting attenuation values for Qs in addition to Vs requires Vp and density information, which are assumed to be available from other geophysical methods, such as seismic reflection (Steeple and Miller, 1990), refraction tomography (Zhang and Toksoz, 1998; Ivanov et al., 2010) for Vp, well measurements, etc. Practical applications of MASW method Vp and density are rarely known with confidence and as a result use of approximations is adopted, which is commonly considered acceptable because of the low sensitivity of surface-wave inversion to Vp and density, i.e., errors in their estimation result in very small errors in final Vs results (Xia et al., 1999). When applied to near-surface seismic data Vp / Vs ratios are rarely below 2.2 or even known. In such cases, when solving for quality factors (Xia et al., 2012) the approximation $Q_p = 2 Q_s$ is assumed.

However, availability of Vp, especially if the corresponding Vp / Vs ratio approaches 2.2 or less makes it possible to invert for Qp in addition to Qs. That is one of the reason for our special interest in ice-sheet data because the reported Vp / Vs ratio for ice can equal 2 (Waite et al., 1995).

Data Acquisition

A 2-D seismic line was acquired at Jakobshavn Glacier, Greenland (69.35° N, 47.2° W), using a 0.5 charge at 10 m depth as a source. The receiver spread consisted of twenty-four 28 Hz vertical phones spaced at 20 m, resulting in a 460 m spread. Data were collected using a roll-along style of acquisition. The spread movement (roll) was 160 m (Horgan et al., 2008).

Results

The MASW technique was appraised by evaluating the ability of the method to estimate the dispersion curve of the fundamental mode of the Rayleigh wave. The phase-velocity – frequency image of data showed a high-resolution fundamental-mode dispersion-curve trend within a frequency range, 10-50 Hz (Figure 1) which could typically be followed to ~110 Hz. With phase velocities ~1600-1700 m/s at ~10 Hz resulted in longest wavelengths of ~160-170 m that contributed to estimating Vs to a depth of ~80 m. A 2-D Vs section was obtained (Figure 2) with the MASW inversion using a Vp model (Figure 3) estimated from a refraction survey in the Antarctic (King and Jarvis, 2007) and density information (Figure 4) from a near-by well. The resultant Vp / Vs ratio was less than 2.2 for the most part of the section (Figure 5).

A set of 18 records was selected to estimate attenuation curves from each record and invert these data for Qs and Qp using the available Vs, Vp and density models. For half of the records the estimated attenuation curves followed close-to-linear trends. For the most of the frequency range (e.g., 20-65 Hz) these curves were matched very well by the curves calculated from the inversion (Figure 6). For a short frequency range (e.g., 10-15 Hz) the curve match was reduced but was considered still acceptable following previous research (Xia et al., 2002; Xia et al., 2012).

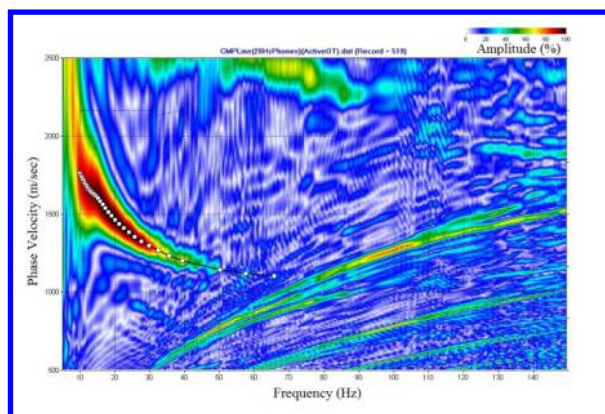


Figure 1. Dispersion-curve images of MASW seismic data acquired at Jakobshavn Glacier, Greenland.

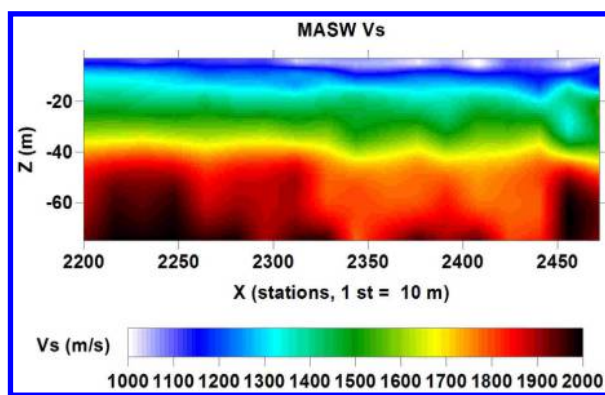


Figure 2. Portion of the 2-D MASW Vs section. Also used as a Vp model with values $V_p = 2V_s$.

For half of the records the attenuation curves appeared irregular (Figure 7), possibly due to noise, and the Q inversion failed to provide a good match between the calculated and estimated curves. These results were not included.

The inverted Qs model (Figure 8) had relatively high values at the very top ~5 m and relatively low Qs for the most of the remaining part of the section (below ~12 m), which implies low shear-wave dissipation at the very top and high dissipation below. The corresponding Qp section (Figure 9) exhibited different trends. There was a relatively high-Qp layer at

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25 m separated by a low-Qp layer at 40 m with two high-Qp dome-shaped anomalies below 45 m.

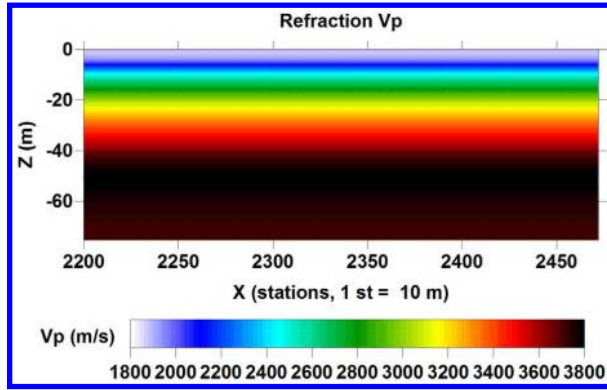


Figure 3. Vp estimated from refraction survey in the Antarctic.

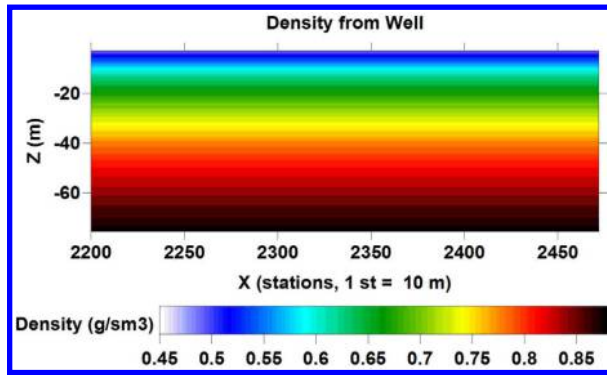


Figure 4. Density information from a near-by well.

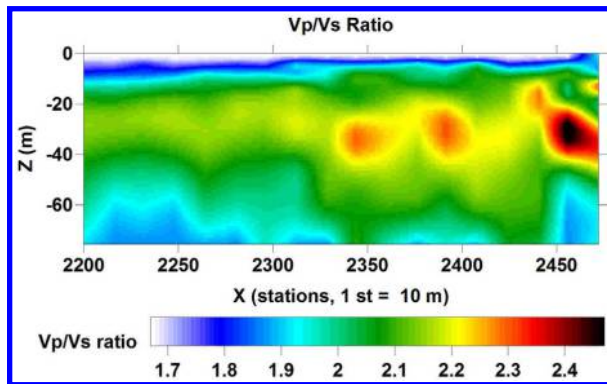


Figure 5. Vp/Vs ratio after inverting for Vs.

To evaluate the impact of the Vp model on the Q inversion, we used a significantly different model using $V_p = 2 V_s$ from MASW (Figure 2; the Vp model would appear identical with values twice bigger) based on Waite et al. (1995). The inverted Qs model (Figure 10) remained almost identical to the

previous with the exception of the increased values of the two previously existing bull-eyes anomalies at about 60 m depth. Qp results changed. The low-Qp layer at 40 m disappeared and the dome-shaped features expanded in width and height (Figure 11). However, both Qp results suggest low dissipation factor (Q_p^{-1}) with depth (high Qp) and high dissipation factor very shallow (low Qp). These relatively high Qp values below ~35 m depth range between 80-100 are consistent with previous glacier research estimates (e.g., 50-160; by using frequencies above 100 Hz) summarized by Gusmeroli et al. (2010). With that respect our estimates may appear a bit lower than the average.

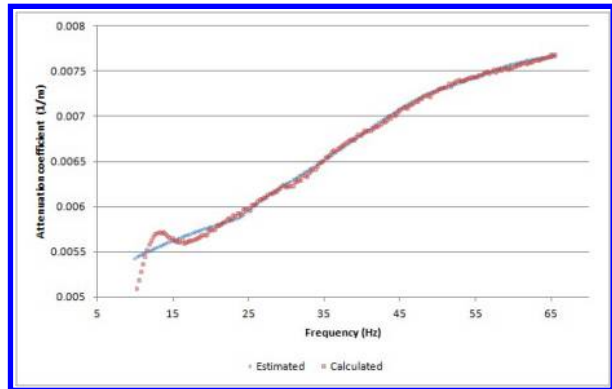


Figure 6. Attenuation curves estimated from the data (blue stars) and calculated (red squares) from the inverted Qs and Qp models.

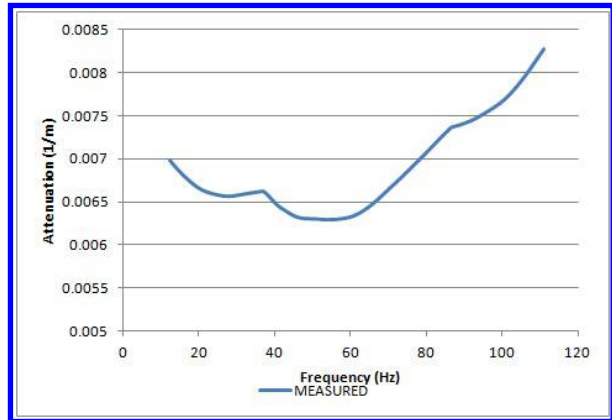


Figure 7. Estimated attenuation curve with irregular trend.

We presented two Vp models (one from well data, and another from lab measurements). Considering the wide possible range of refraction nonuniqueness, more Vp models can be obtained using algorithms, such as, GRM (Palmer, 1981), refraction tomography (Zhang and Toksoz, 1998), JARS (Ivanov et al., 2010), etc. for potentially detailed Qp results (e.g., Figure 9).

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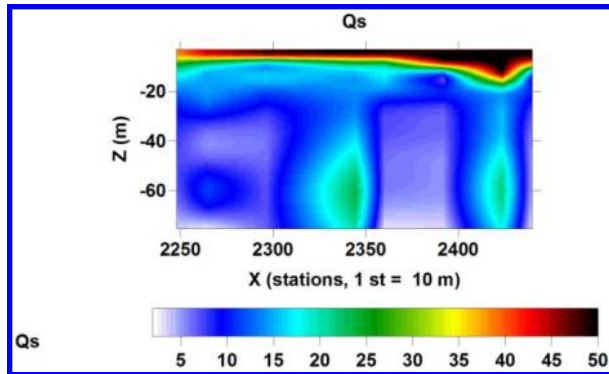


Figure 8. Q_s estimates using MASW Vs and refraction V_p .

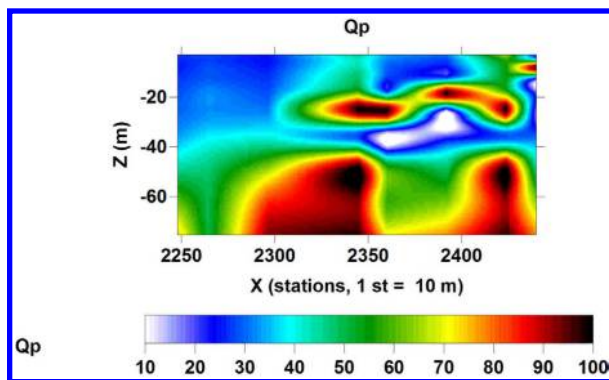


Figure 9. Q_p estimates using MASW Vs and refraction V_p .

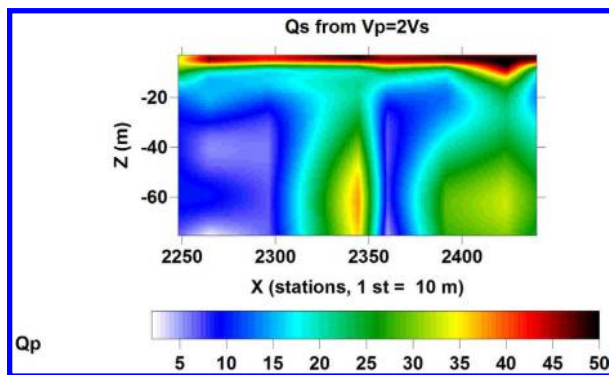


Figure 10. Q_s estimates using MASW Vs and $V_p=2V_s$.

This research showed that using the significantly different V_p models resulted in different Q_p ; however, the overall Q_p trend is preserved. As well, Q_s is largely unaffected by changes in the V_p model.

High Q_s and Q_p (lower dissipation factors Q_s^{-1} and Q_p^{-1}) with depth are consistent with our expectations of glacier changes with depth.

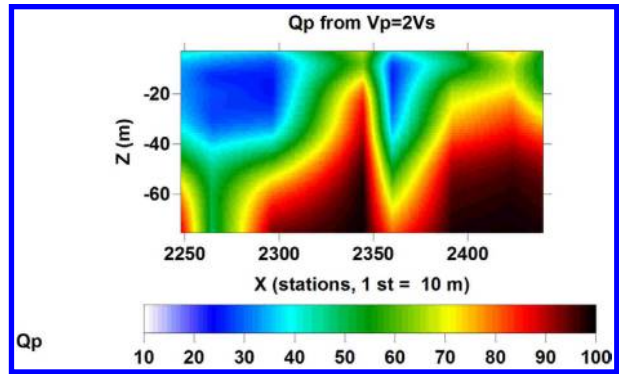


Figure 11. Q_p estimates using MASW Vs and $V_p=2V_s$.

Conclusions

This work shows that in addition to estimating quality factor Q_s from Rayleigh wave attenuation coefficients, it is also possible to obtain from these same data Q_p estimates that look different from the Q_s trends (i.e., they were not a scaled version of Q_s). This was possible only by using V_p models that provide V_p/V_s ratio < 2.2 .

To the best of our knowledge our work is a unique attempt to estimate Q_p from Rayleigh-wave attenuation coefficients for the very near-surface (i.e., 10-80 m), for which V_p/V_s ratio are < 2.2 . We speculate that this may be due to the fact that at great majority of the very near-surface sites $V_p/V_s > 2$.

Future research can include the estimation of more accurate V_p models for this site using other geophysical methods, such as refractions, tomography, well measurements, etc.

We demonstrate that the seismic method can provide more than one physical attribute for observing status and processes taking place in a glacier. In this instance a single seismic survey for reflection purposes could also be used for the estimation of 2-D MASW Vs, Q_s , and Q_p sections.

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EDITED REFERENCES

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